

# Mechanisms of Salinity Anomalies in the Bohai Sea in January 2007

Fengying Ji<sup>1</sup> and Xuejun Xiong<sup>2</sup>

<sup>1</sup>National Marine Data and Information Service China

<sup>2</sup>First Institute of Oceanography, Ministry of Natural Resources

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## Abstract

In view of the fact that the practical salinity of the Bohai Sea in January 2007 was significantly higher than the multi-year mean with its horizontal distribution opposite to the latter, the factors that affect the interannual variation of practical salinity in the Bohai Sea are quantitatively analyzed based on the in-situ hydrological measurements, the annual runoff data of the Yellow River into the sea, as well as the precipitation and evaporation reanalysis data of EAR5. The results show that the local freshwater supply not only dominates the magnitude of salinity change in the Bohai Sea, but also causes the salinity in the central Bohai Sea is higher than that in the Bohai Strait in winter in some years which is in inverse to the climatological salinity field. The seesaw distribution characteristics of the freshwater supply of the Bohai, Yellow and East China Seas also strengthen the characteristics that the salinity horizontal distribution opposite to the climatological in the Bohai Sea in some years. Available observations also show that the nutrient and inorganic carbon of the Bohai Sea are much higher than that of the open ocean, which gives a rise of  $0.02 \sim 0.2$  g·kg<sup>-1</sup> in Absolute Salinity. Therefore, it is necessary to replace the Practical Salinity with the Absolute Salinity for the accurately salinity changes study in the Bohai Sea.

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# Mechanisms of Salinity Anomalies in the Bohai Sea in January 2007

Ji Fengying<sup>1</sup>, Xiong Xuejun<sup>2</sup>

<sup>1</sup>National Marine data and information service, Tianjin, 300171, China.

<sup>2</sup> First institute of oceanography, ministry of natural resources, Qingdao, 266001, China.

Corresponding author: Fengying Ji ([2320130582@qq.com](mailto:2320130582@qq.com))

## Key Points:

- The local freshwater supply dominates the salinity change of the Bohai Sea.
- The freshwater supply in the Bohai, Yellow and East Seas is characterized by a seesaw distribution, which strengthens the anti-phase distribution of Bohai Sea.
- The relative composition of the Bohai Seawater is different from the open ocean which has an impact on salinity changes.

## Abstract

In view of the fact that the practical salinity of the Bohai Sea in January 2007 was significantly higher than the multi-year mean with its horizontal distribution opposite to the latter, the factors that affect the interannual variation of practical salinity in the Bohai Sea are quantitatively analyzed based on the in-situ hydrological measurements over 35 years, the annual runoff data of the Yellow River into the sea, as well as the precipitation and evaporation reanalysis data of EAR5. The results show that the local freshwater supply not only dominates the magnitude of salinity change in the Bohai Sea, but also causes the salinity in the central Bohai Sea is higher than that in the Bohai Strait in winter in some years which is in inverse to the climatological salinity field. The seesaw distribution characteristics of the freshwater supply of the Bohai, Yellow and East China Seas also strengthen the characteristics that the salinity horizontal distribution opposite to the climatological in the Bohai Sea in some years. Available observations also show that the nutrient and inorganic carbon of the Bohai Sea are much higher than that of the open ocean, which gives a rise of  $0.02 \sim 0.2 \text{ g} \cdot \text{kg}^{-1}$  in Absolute Salinity. Therefore, it is necessary to replace the Practical Salinity with the Absolute Salinity for the accurately salinity changes study in the Bohai Sea.

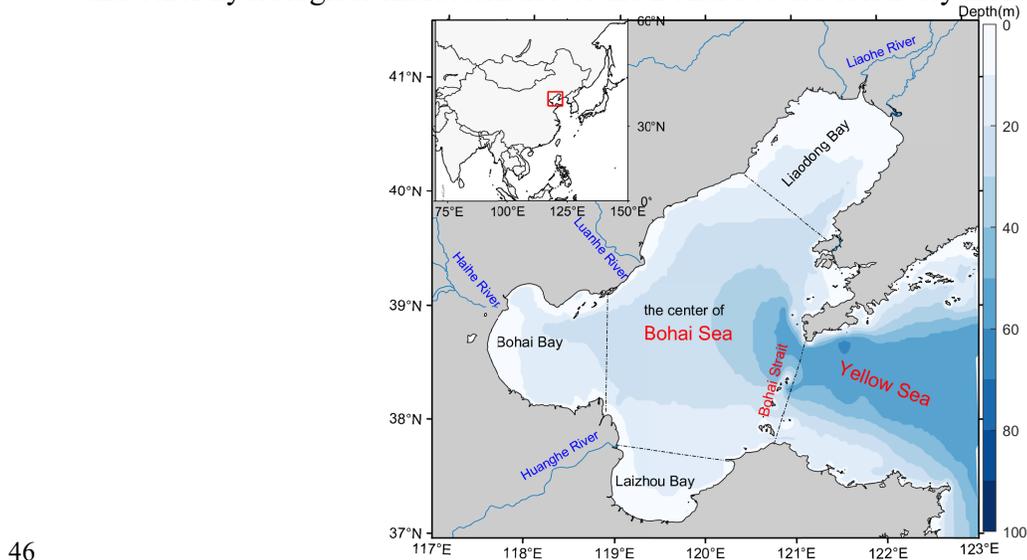
## Plain Language Summary

Based on a large amount of in-situ hydrographic measurements and reanalysis data, this paper reveals the mechanism of that the practical salinity of the Bohai Sea in January 2007 was not only significantly higher than the multi-year mean, but also the horizontal distribution was opposite to the latter, as well as indicates it is necessary to replace the Practical Salinity with the Absolute Salinity for accurately study of salinity changes of the Bohai Seawater.

## 1 Introduction

The Bohai Sea is the largest inland sea in China, surrounded by land on three sides and connected to the Yellow Sea only through the Bohai Strait, which is only 105 km wide in the southeast, and is a typical semi-enclosed bay with a small mouth and a large belly. The central

39 basin of the Bohai Sea, facing the Bohai Strait directly, is surrounded by three shallow terrace-  
 40 type bays, Liaodong Bay in the northeast, Bohai Bay in the west and Laizhou Bay in the south.  
 41 The average depth of the Bohai Sea is only 18 meters, the deepest depth is 86 meters, located in  
 42 the northern end of the Bohai Strait in the Old Tieshan waterway. Because of the shallow water  
 43 and strong confinement, the Bohai Sea is significantly influenced by the continental climate;  
 44 coupled with the influence of river runoff injected into it, the temperature and salt distribution  
 45 and other hydrological characteristics of the Bohai Sea are relatively isolated and variable.



46

47 **Figure 1.** The Bathymetry of the Central Bohai Sea, data from Etopo 2

48 It's found that the salinity of the Bohai Sea in January 2007 acquired from a special  
 49 project entitled Marine Integrated Investigation and Evaluation Project of China Offshore  
 50 conducted from 2006 to 2007 (hereafter referred to as the '06's Investigation') is not only larger  
 51 than multi-year averaged winter salinity(Chen, 1992; SOA, 2016; Zweng et al, 2019) , but also  
 52 the horizontal distribution was opposite to the latter: the climate atlas shows that the salinity of  
 53 Bohai Sea decreases from the eastern Bohai Strait mouth gate to the north, west and south,  
 54 whereas the salinity inside the Bohai Sea is significantly higher than that in the Bohai Strait in  
 55 January 2007, with the maximum value occurring along the northwest coast, and then gradually  
 56 decreasing towards the eastern Bohai Strait mouth gate.

57 Salinity is a term used to quantify the total mass of inorganic substance dissolved in pure  
 58 water to form a given mass of seawater which is as a conserved seawater property (IOC et al,  
 59 2010), now classified as an 'Essential Climate Variable' (GCOS, 2010). Local long-term trends  
 60 in salinity are precisely measurable indicators for climatic changes in the terrestrial water cycle  
 61 and its sensitivity to global warming (Pawlowicz et al, 2016). Such a significant change has  
 62 attracted many oceanographers to conduct related research.

63 Based on observations of the fixed stations along the coast of Bohai Sea and the section  
 64 in the central of the Bohai Sea, it was found that the salinity of seawater in the Bohai Sea has  
 65 been increasing continuously since 1960, and the reasons for this are the sharp decrease in the  
 66 runoff of the Yellow River, the change in the intensity of the invasion of the Yellow Sea Warm  
 67 Current, and the influence of precipitation and evaporation (Lin et al., 2001; Fang et al. ,2002;  
 68 Wu et al., 2004a, 2004b; Zu et al, 2005; Xu et al, 2007; Ma et al., 2006, 2010; Jia et al., 2008);

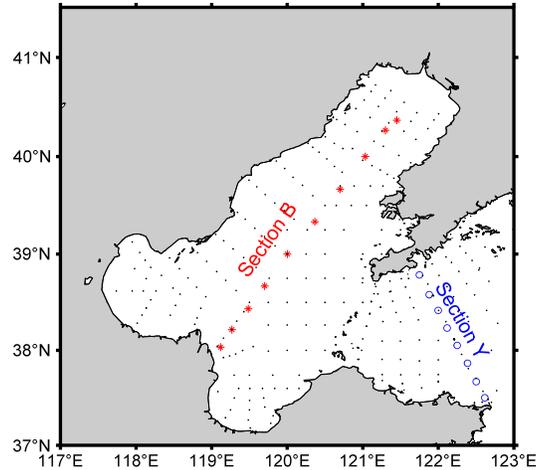
69 Based on the results of the interannual statistical analysis of the salinity data of the Bohai Sea  
70 from 1950 to 1990 based on field observations, it is not only found that the salinity of the whole  
71 Bohai Sea is increasing, but also the salinity horizontal distribution of the Bohai Sea has changed  
72 fundamentally since the 1980s, and Bohai Bay has become the high value area of the Bohai Sea  
73 salinity (Yu et al., 2015); based on quasi-synoptic special investigation data, it was also found  
74 that the salinity in the Bohai Sea was significantly higher in summer 2000 and 2008 than that in  
75 summer 1958 (Wu et al, 2004b; Song et al., 2009).

76 The above studies obtained the magnitude and rate of salinity increase in the Bohai Sea  
77 over past decades by linear fitting, and the causes of this phenomenon by correlation analysis,  
78 but did not quantitatively analyze the changing patterns and influencing factors of Bohai Sea  
79 salinity on an annual scale; moreover, they did not explain the higher salinity in the central of  
80 Bohai Sea than in the Bohai Strait in January 2007 which is inverted to the climatological atlas.  
81 In this paper, we firstly use the in-situ observations, including quasi-synchronous field  
82 investigation, 23 years of repeat hydrographic measurements along the section B in the central  
83 Bohai Sea and the section Y crossing the Bohai Strait, and the runoff data of the Yellow River  
84 into the sea, and precipitation and evaporation reanalysis data of EAR5 to quantitatively  
85 calculate the influences of annual freshwater budget on the salinity in the Bohai Sea. Due to the  
86 poor water exchange between the Bohai Sea and the open sea, the local freshwater budget not  
87 only dominate the salinity magnitude and distribution, but also lead to the seawater nutrient and  
88 inorganic carbon significantly higher than the open ocean, resulting in the Absolute Salinity will  
89 increase by  $0.02 \sim 0.2 \text{ g}\cdot\text{kg}^{-1}$ , which cannot be ignored in the salinity study. Therefore, in order to  
90 accurately study the salinity variation in the Bohai Sea, Absolute Salinity should be used instead  
91 of Practical Salinity.

## 92 **2 data**

### 93 2.1 '06's Investigation' data

94 The winter salinity observation in the Bohai Sea area was from December 28, 2006 to  
95 January 23, 2007, obtained by the CTD produced by Alec Company of Japan (Xiong, 2012). The  
96 seawater salinity was calculated by temperature, conductivity, Pressure by EOS-80, the stations  
97 location is shown in Fig.2.



98

99 **Figure 1.** The location of the observations of '06's Investigation(.), Section B(\* ) and 'Section  
100 Y (° )

### 101 2.2 Repeat hydrographic measurements along sections in the Bohai Sea over 35 years

102 The station locations and observation time of the repeat hydrographic measurements  
103 along sections in the Bohai Sea have stabilized since 1986. Section B, which runs through the  
104 central Bohai Sea, and Section Y, which crosses the Bohai Strait, with station locations shown in  
105 Fig. 2, basically are ensured to be investigated twice a year, with winter and summer  
106 observations made in February and August each year.

107 The data of sections in the winter of 1996 and 1999 when no observations were carried,  
108 are derived by linear interpolation based on the data of the two years before and after. The  
109 observation of the missing stations in 1987, 1988, 1998, 2001 and 2002, are obtained by linear  
110 interpolation of the data of the same level of the adjacent stations. Thus, a continuous series of  
111 the measurements along the section at fixed stations and fixed observation levels is formed,  
112 which ensures the reliability and comparability of the average salinity of the section.

### 113 2.3 ERA5 data

114 ERA5 is the fifth generation ECMWF atmospheric reanalysis of the global climate  
115 covering the period from January 1950 to present. ERA5 is produced by the Copernicus Climate  
116 Change Service (C3S) at ECMWF which could be download: <https://cds.climate.copernicus.eu>.

### 117 2.4 Runoff data of the Yellow River into the Bohai Sea

118 The runoff data of the Yellow River into the Bohai Sea from 1998 to 2020 comes from  
119 the Yellow River Water Resources Bulletin released by Yellow River Conservancy Commission  
120 of Ministry of Water Resource(hereafter referred to as YRCC/MWR), which can be downloaded:  
121 [www.yrcc.gov.cn/zwzc/gzgb/gb/szygb](http://www.yrcc.gov.cn/zwzc/gzgb/gb/szygb).

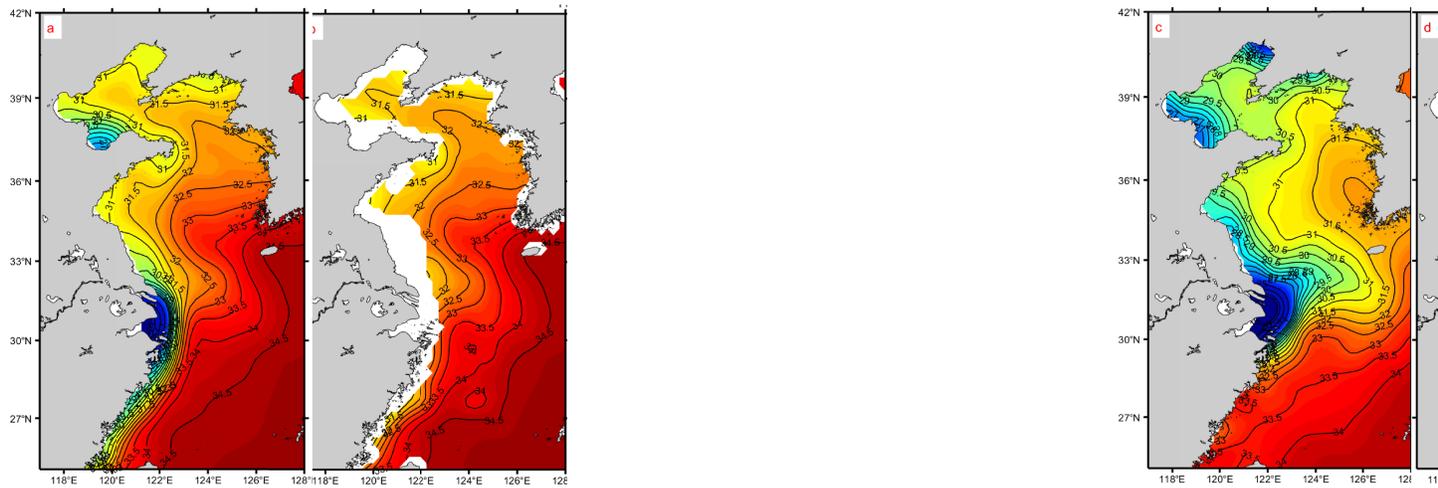
### 122 3 Salinity field in the Bohai Sea

#### 123 3.1 Climatological salinity field in the Bohai Sea

124 The multi-year averaged salinity field in the Bohai Sea is analyzed based on the China  
125 Offshore Marine Atlas (Chen, 1992; SOA, 2016; Zweng et al, 2019), as shown in Fig. 3.

126 In winter, under the combined effect of prevailing strong northerly monsoon and sea  
127 surface cooling, the vertical convection and eddy mixing of seawater is very strong, and the  
128 spatial distribution of salinity of the water mass under the surface is basically the same as that of  
129 the surface in the Bohai Sea. In winter, the SSS (sea surface salinity) in the Bohai Sea has a  
130 strong continuity, and a salinity tongue as high as 31.5 from the northern Yellow Sea through the  
131 northern part of the Bohai Strait to the center of the Bohai Sea, so that the salinity decreases from  
132 the center to the north, west and south in three directions.

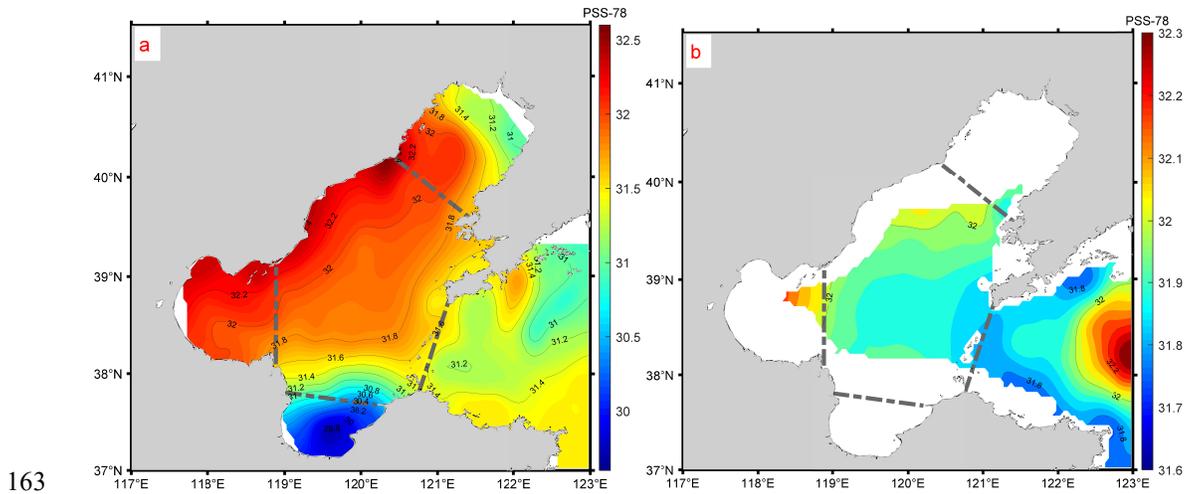
133 In summer, the south wind prevails in the Bohai Sea, precipitation is concentrated and  
134 reaches the maximum in a year, river runoff into the sea is the largest, the SSS drops to the  
135 lowest throughout the year, and SSS decreases horizontally from the central Bohai Strait to the  
136 north, west and south in three directions, the lowest in the top of the three bays, as shown in Fig.  
137 3 c. Due to the smaller wind speed and the strongest solar radiation in summer, the sea surface  
138 warms up rapidly, seawater stability increases, eddy mixing becomes more difficult, and  
139 seawater salinity stratification is obvious. The salinity tends to increase below the sea surface. In  
140 the 20 m depth, the central of the Bohai Sea is mostly controlled by saltier water above 30.5, as  
141 shown in Fig.3 d.



158 **Figure 2.** Salinity isoclines at (a) sea surface in winter (b) 20 m depth in winter (c) sea surface in  
159 summer (d) 20 m depth in summer

#### 160 3.2 Salinity field in the Bohai Sea in winter 2007

161 Based on the observation of the '06's Investigation', the salinity isolines on different  
162 depths in the Bohai Sea in winter 2007 are shown in Fig. 4.



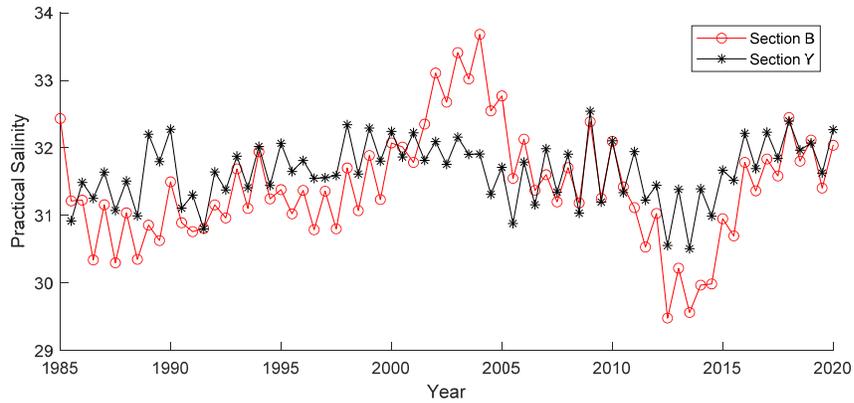
163  
 164 **Figure 3.** The winter salinity isoclines at (a) sea surface (b) 20 m depth of the Bohai Sea in '06's  
 165 Investigation'

166 In winter, the highest SSS of the Bohai Sea, up to 32.45, is along the northwest coast, and  
 167 the lowest, as low as 30.6, is in the southeast near Laizhou Bay due to its proximity to the mouth  
 168 of the Yellow River. The saltier tongue can be divided into three parts, the northwest end near  
 169 the coast has the most drastic change in salinity, from 32.6 to 31.9. while the salt variation in the  
 170 center of the Bohai Central Sea is very smooth, ranging from 31.8 to 31.9, while at the mouth of  
 171 the Bohai Strait the change is stronger, from 31.8 to 31.4 as shown in Fig.4 a.

172 The salinity at 20 m in the Bohai Sea does not vary much, ranging from 31.9 to 32, and is  
 173 still higher in the northwest than in the Bohai Strait, which is in the southeast, like the SSS  
 174 distribution.

### 175 3.3 Time series of averaged salinity of section B and section Y

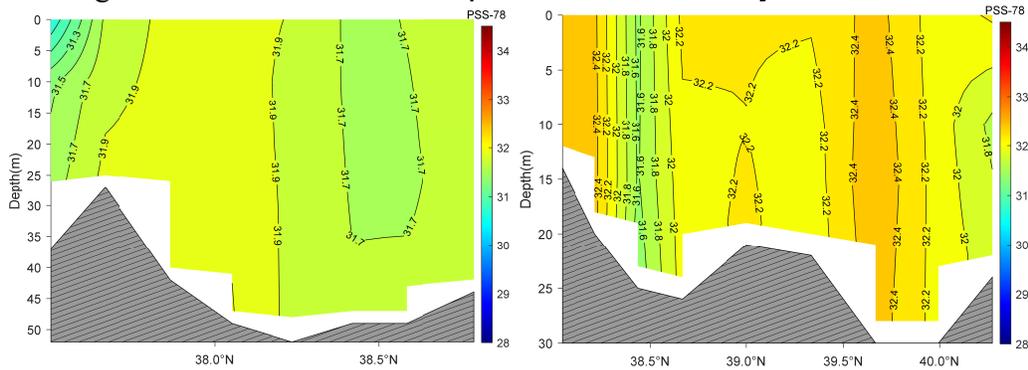
176 As shown in Fig. 5, obvious seasonal salinity variations both appear in section B and Y  
 177 during the existing 35-year observation serials. The mean salinity of summer section is  
 178 significantly lower than that of winter section in the same year; the seasonal and annual  
 179 variations of salinity of section B are significantly higher than that of section Y; excluding  
 180 August 2001 to August 2006, winter 2018 and winter 2019, the salinity of section Y is higher  
 181 than that of section B in the remaining 28 years. The salinity of section B started to be higher  
 182 than that of section Y in August 2001, and then increased continuously, reaching a maximum  
 183 value of 33.5 in 2003, which was 1.6 above the average salinity of section Y. The salinity of  
 184 section B then decreased rapidly and was already significantly lower than that of section Y in  
 185 February 2007, and continued until 2017. Subsequently, the phenomenon of inverse salinity  
 186 distribution occurred again in 2018 and 2019.



187

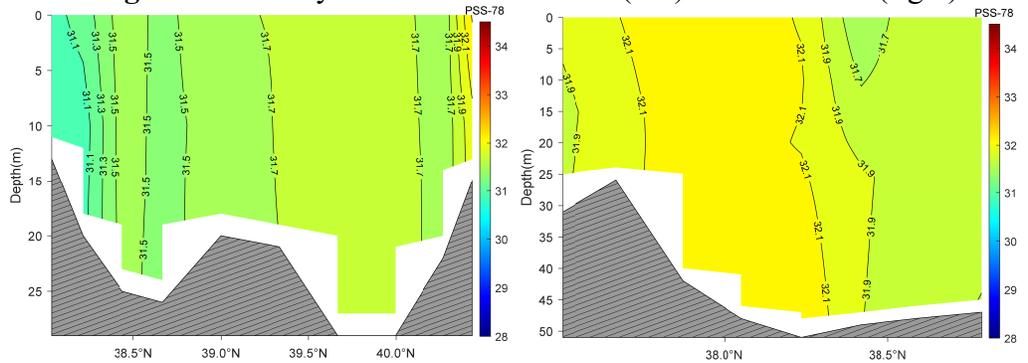
188 **Figure 4.** The average salinity of section B (red o) and Y (black \*) from 1985 to 2020

189 Since the winter investigation period of ‘06’s Investigation’ was conducted in the Bohai  
 190 Sea from December 2006 to January 2007, which was located between the summer 2006 and  
 191 winter 2007 observations of the section, it can be seen in the Fig.6 and Fig.7 that the salinity of  
 192 section B in August 2006 was significantly higher than that of section Y, but by February 2007,  
 193 the average salinity of section B was already slightly lower than that of section Y. Although the  
 194 ocean dynamics processes occurring in between are not clear, it still indicates that the ‘06’s  
 195 Investigation’ is consistent with the phenomena reflected by the cross-sectional observations.



196

197 **Figure 5.** Salinity isoclines of section B (left) and section Y (right) in August 2006



198

199 **Figure 6.** Salinity isoclines of section B (left) and section Y (right) in February 2007

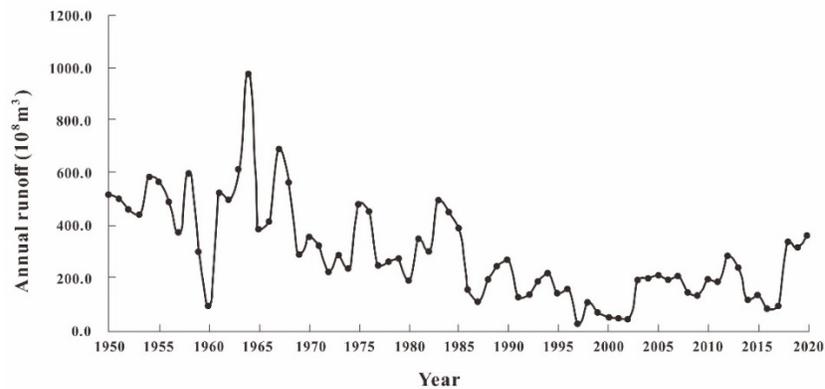
#### 200 **4 Mechanisms of salinity anomalies in the Bohai Sea in winter 2007**

201 According section 3, in most cases, the salinity in the Bohai Sea in winter is usually  
 202 lower than that in the Bohai Strait, but in some years this distribution is reversed. As a measure  
 203 of dissolved substances in seawater, the variation of salinity in the Bohai Sea is the result of local  
 204 freshwater supplies, the replenishment of external seawater and the ionic substance input.  
 205 Therefore, this section will begin with above factors.

##### 206 4.1 Local freshwater budget of the Bohai Sea

207 The total area of the Bohai Sea is  $7.73 \times 10^{10} \text{ m}^2$ , with an average depth of 18 m and a total  
 208 storage capacity of  $1.39 \times 10^{12} \text{ m}^3$ .

209 There are many rivers discharge into the Bohai Sea along the coast, but the runoff is  
 210 much smaller than that of the Yellow River. Due to the climate change and human activities, the  
 211 runoff discharging into the Bohai sea of the Yellow River which can be basically expressed by  
 212 the measurements of its most downstream Lijin hydrological station vary significantly, as shown  
 213 in Fig.8, with an average annual runoff  $3.1319 \times 10^{10} \text{ m}^3 \cdot \text{a}^{-1}$  from 1956 to 2000, and  
 214  $1.39 \times 10^{10} \text{ m}^3 \cdot \text{a}^{-1}$  from 1987 to 2000 (YRCC/MWR, 2020).

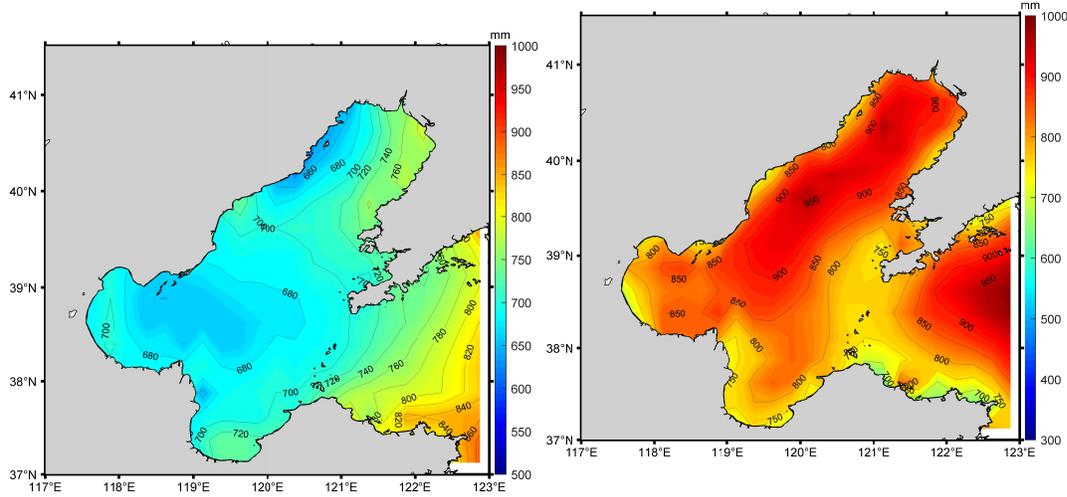


215

216 **Figure 7.** The annual runoff of the Yellow River from 1950 to 2020 of Lijin Hydrological  
 217 Station

218 Due to the lack of in-situ measurement series with wide coverage and long-time span, the  
 219 precipitation data of ERA5 in the Bohai Sea are evaluated with the present limited measurement  
 220 of stations along the coast firstly. The multi-year averaged annual precipitation in the Bohai Sea  
 221 from 1960 to 2018 is  $536 \text{ mm} \cdot \text{kg}^{-1} \cdot \text{a}^{-1}$  based on the measurements of coastal stations (Fan et al,  
 222 2021) and is  $710.27 \text{ mm} \cdot \text{kg}^{-1} \cdot \text{a}^{-1}$  according to the ERA5 whose distribution is higher in the  
 223 central of the Bohai Sea than along the coast, shown in Fig. 9 (left). All above indicates that the  
 224 ERA5 precipitation data in the Bohai Sea makes no difference with the measurements on the  
 225 climatic scale and can be applied in the analysis.

226 The multi-year averaged annual evaporation of ERA5 is about  $1170.92 \text{ mm} \cdot \text{kg}^{-1} \cdot \text{a}^{-1}$  from  
 227 1960 to 2018, which is 1.65 times the multi-year averaged annual precipitation, with high values  
 228 in the central and northern Bohai Sea as shown in Fig. 9 (right), which indicates that the Bohai  
 229 Sea, located in the mid-latitude monsoon region, has a dry climate.



230 **Figure 8.** The isolines of multi-year averaged annual precipitation (left) and evaporation (right)  
 231 of EAR 5 from 1960 to 2018

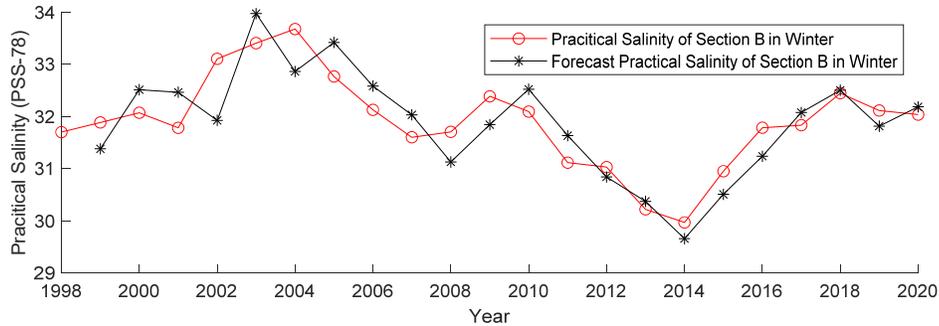
232 Ignoring the weak water exchange of Bohai Sea with the open sea first, the multi-year  
 233 averaged annual freshwater budget (the sum of precipitation and runoff minus evaporation) is 2  
 234 orders of magnitude smaller than the total water storage in the Bohai Sea due to the annual  
 235 runoff of the Yellow River into the Bohai Sea of  $2.72 \times 10^{10} \text{ m}^3 \cdot \text{a}^{-1}$ , the annual evaporation of  
 236  $9.0 \times 10^{10} \text{ m}^3 \cdot \text{a}^{-1}$  and the annual precipitation of  $5.5 \times 10^{10} \text{ m}^3 \cdot \text{a}^{-1}$  from 1959 to 2020. The following  
 237 approximate formula is used to calculate the annual change in salinity of the Bohai Sea due to  
 238 freshwater budget according to the salinity definition,

$$239 \quad \Delta \text{Psal}_i = \text{Psal}_i \times \text{WaterStorage} / (\text{freshWaterBudget}_i + \text{WaterStorage}) - \text{Psal}_i \quad (1)$$

240 In the above formula,  $\Delta \text{Psal}_i$  is the salinity increment of the  $i+1$ th year relative to the  $i$ th  
 241 year due to the freshwater budget,  $\text{Psal}_i$  is the average salinity of the Bohai Sea in the  $i$ th year,  
 242 replaced by the mean salinity of section B in winter of the  $i$ th year,  $\text{freshWaterBudget}_i$  is the  
 243 fresh water supply of the Bohai Sea in the  $i$ th year, which is the annual runoff of the Yellow  
 244 River, represented by  $\text{YellowRiver\_Runoff}_i$ , and the annual precipitation denoted  
 245  $\text{Bohai\_Precipitation}_i$  minus the annual evaporation of the Bohai Sea, denoted  
 246  $\text{Bohai\_Evapouration}_i$  in the Bohai Sea in the  $i$ th year, that is,

$$247 \quad \text{freshWaterBudget}_i = \text{YellowRiver\_Runoff}_i + \text{Bohai\_Precipitation}_i - \text{Bohai\_Evapouration}_i \quad (2)$$

248 Since only the annual runoff of the Yellow River after 1998 is released, only the  
 249 correlation between the mean salinity of Section B in winter and the predicted salinity values  
 250 based on the freshwater budget and the salinity mean of Section B in the previous year from  
 251 1999 to 2020 is calculated. The correlation between observations and the predictions is 0.85 with  
 252 the confidence level of 0.95, and the average of salinity difference is 0.4, as shown in Fig.10.



253  
254 **Figure 9.** The mean salinity ( $^{\circ}$ ) of Section B in winter and the predicted salinity based on  
255 the freshwater budget of the Bohai Sea (\*)

256 As shown in Fig.10, the freshwater supply of the Bohai Sea derived by the Yellow River  
257 runoff, evaporation, and precipitation in the previous year basically dominate the salinity change  
258 in the subsequent year, and also explains why the salinity of the central of Bohai Sea was higher  
259 than that of the Bohai Strait in 2001-2006, 2017, and 2018: the freshwater supply of the Bohai  
260 Sea continued to be negative from 2000 to 2002, resulting in the salinity of the Bohai Sea keep  
261 rising; although the freshwater budget of the Bohai Sea was positive from 2003, the net increase  
262 of freshwater supply was not enough to reduce the salinity of the central Bohai Sea to that of  
263 Bohai Strait until 2007 when the horizontal distribution of salinity return to the climate state  
264 again. In 2016 and 2017, along with the rising of the evaporation and decreased precipitation, the  
265 Yellow River runoff in this two years was lower than in previous years again, as shown in Fig.8,  
266 resulting in higher salinity in the central of Bohai Sea than that of the Bohai Strait in 2018 and  
267 2019.

268 Since the average difference between the predicted salinity and the measured salinity is  
269 as high as 0.4 this algorithm is useful but relatively coarse, and there should be other factors to  
270 affect the salinity change in the Bohai Sea, such as the water exchange between the Bohai Sea  
271 and the Yellow Sea.

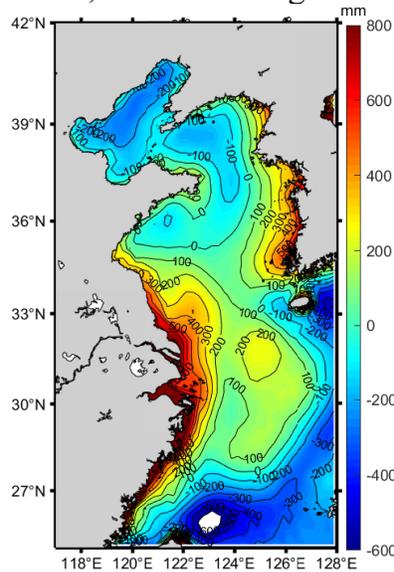
#### 272 4.2 Freshwater budget of the Bohai, Yellow and East China Seas

273 The seawater of the Bohai, Yellow and East China Seas could be roughly considered as a  
274 mixture of two water masses, one is the fresher water mass along China coast formed by the  
275 runoff of rivers into the sea and the other is saltier water mass of the China offshore brought by  
276 the Kuroshio Current and its branches. The distribution characteristics of the salinity of Chinese  
277 offshore seawater are characterized by: fresher on the surface and saltier in the deep layer;  
278 fresher nearshore and saltier offshore; freshest in the estuarine and saltiest in the Kuroshio, as  
279 shown in Fig.3.

280 As the second strongest current in the world after the Gulf Stream, the Kuroshio  
281 originates from the northward branch of the North Equatorial Current in the Pacific Ocean after  
282 it reaches Luzon, acts as a natural barrier between North West Pacific seawater and the East  
283 China Sea seawater. Compared with the circulation in the Chinese shelf seas, the Kuroshio  
284 Current is more stable, more continuous and less seasonal variation. Therefore, this paper ignores  
285 the impact of its salinity variability on the Chinese offshore seawater.

286 As the largest river in China, the Yangtze River is the main body of China's coastal  
 287 diluted water and its annual runoff discharging into the sea is maintained at 1012 m<sup>3</sup> which is the  
 288 same order of magnitude as the total water storage capacity of the Bohai Sea. In the flood season  
 289 (summer), the Yangtze River's diluted water is forced by the southerly wind and pushed by the  
 290 Taiwan Warm Current to the north, its water body shifts northward and most of water discharged  
 291 to the sea are held near the Yangtze River estuary, as shown in Fig.3 c, d. In autumn and winter,  
 292 the north wind strengthens, the solar radiation weakens, the cooling speeds up, the convection  
 293 and eddy mixing strengthen, and the fresh water on the ocean surface is transported to the  
 294 bottom, maintain the salinity of the Bohai, Yellow and East China Seas is significantly lower  
 295 than that of Kuroshio Current and the ocean west of it, shown in Fig.3 a, b (Xiong, 2012).

296 Based on the ERA5 from 1959 to 2021, the annual average of freshwater budget of China  
 297 offshore is negative north of 35°N and is positive in a fan-shaped area centered on the Yangtze  
 298 River, from 34°N in the north to 27°N in the south and 127°E in the west, the maximum locates  
 299 in the estuary of Yangtze River. Its isolines are in good agreement with the salinity contours  
 300 Outside the mouth of the Yangtze River, as shown in Fig.11 and Fig.3.



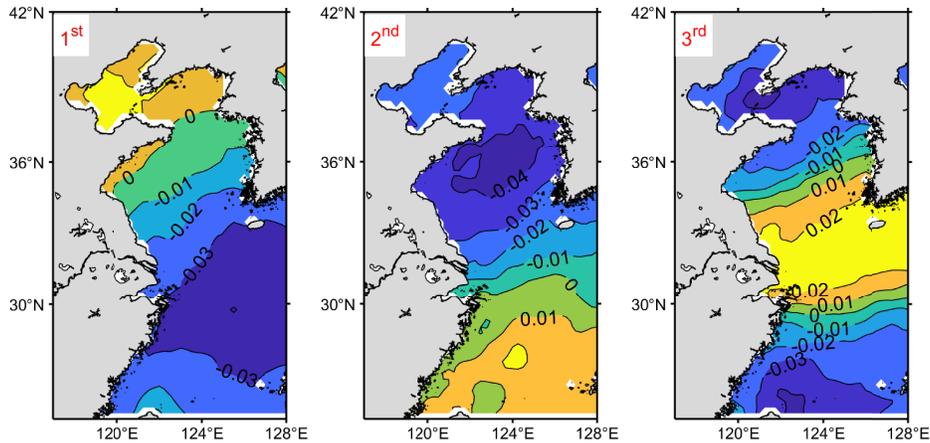
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 302

**Figure 10.** Annual average of freshwater budget from 1959 to 2020

303 To further understand the spatial and temporal characteristics of the interannual variation  
 304 of freshwater budget for the Bohai, Yellow and East China Seas, 63 years of monthly average  
 305 freshwater budget series from 1959 to 2021 were subjected to 12 months of low-pass filtering (to  
 306 eliminate the effect of seasonal variation), and then the obtained annual components were  
 307 subjected to an empirical orthogonal function (EOF for short) decomposition.

308 **Table 1** The Eigenvalues and Percentages of the First Three Eigenvectors of Freshwater Budget  
 309 of the Bohai, Yellow and East China Seas

Mode	1 <sup>st</sup>	2 <sup>nd</sup>	3 <sup>rd</sup>	Sum
Eigen value	475.4	390.8	239.5	1105.8
Percentage /%	28.52	23.45	14.37	66.33

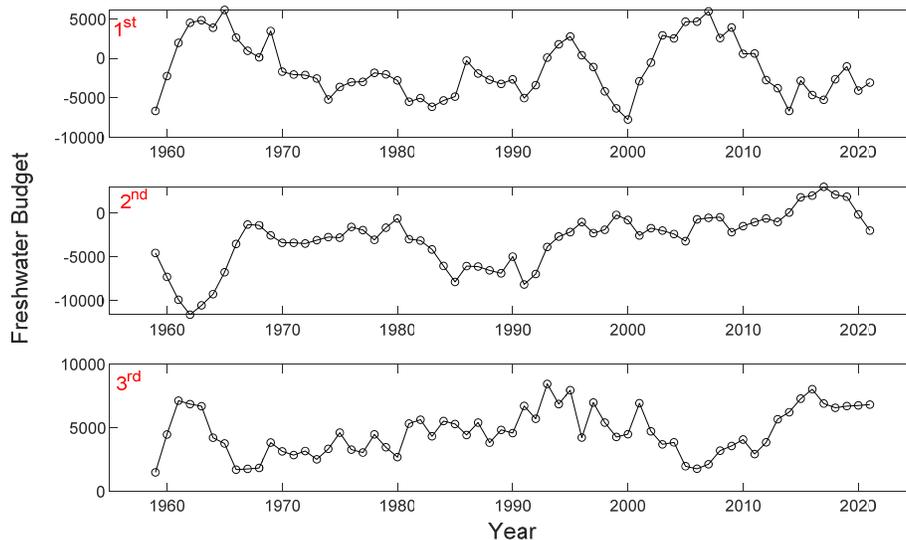


310

311

**Figure 11.** The first three spatial patterns of EOFs of freshwater budget from 1959 to 2021 in Bohai Sea, Yellow Sea and East China Sea.

312



313

314

**Figure 12.** The first time coefficients of three temporal EOFs of freshwater budget from 1959 to 2021 in Bohai Sea, Yellow Sea and East China Sea.

315

316 As shown in Table 1 and Fig.12, the first three spatial EOFs are basically distributed  
 317 along the latitude, showing a "seesaw" structure of north-south inverse phase oscillation, which  
 318 is also the expression of the balance of freshwater in China's offshore, with an increase in the  
 319 south and a corresponding decrease in the north, and vice versa. The first EOF explains 27.73%  
 320 of the total variance, which is the main pattern of freshwater budget in the Bohai Sea, the Yellow  
 321 and East China Seas. It is bounded by 34°N~36°N, with negative values south of the boundary  
 322 and positive values north of it, indicating that when freshwater budget is increasing south of the  
 323 boundary, it decreases north of the boundary. The coefficients of their temporal changes from  
 324 1998 to 2002, and from 2017 to 2019 are negative, shown in Fig.13, indicating that the  
 325 freshwater balance of the northern Bohai Sea is negative and the freshwater balance of the  
 326 southern Yellow Sea and the East China Sea is positive, thus leading to the increasing salinity of  
 327 the Bohai Sea in these years and the decreasing salinity of its outer oceanic water, together with

328 the insufficient exchange of seawater in the Bohai Sea, the cumulative effect leads to the increase  
 329 of the salinity inside the Bohai Sea and causes this inverse phase of the salinity inside the Bohai  
 330 Sea being higher than the salinity of the Yellow Sea.

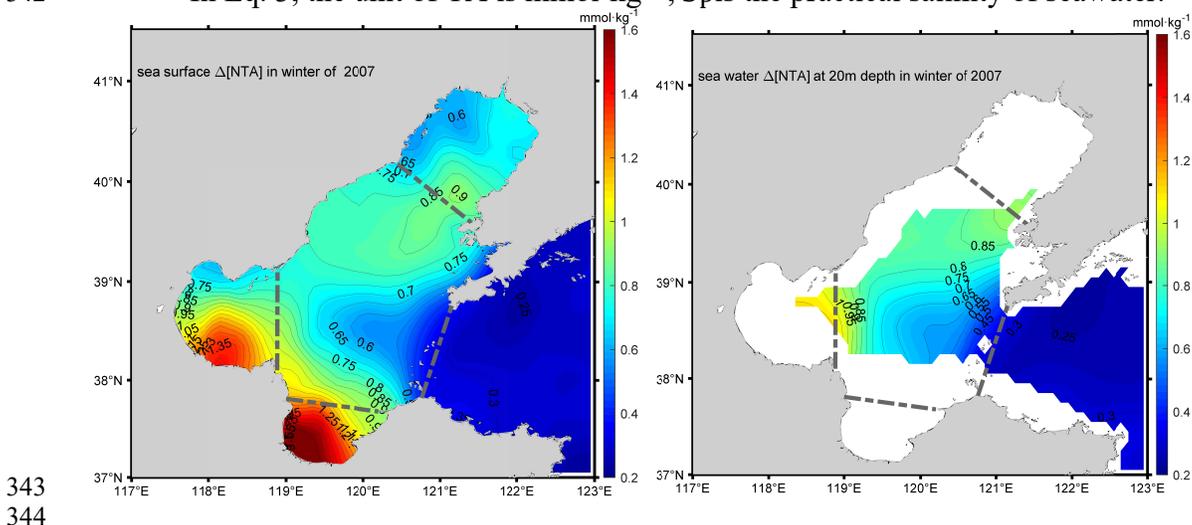
### 331 4.3 Local ionic substance input to the Bohai seawater

332 From Section 4.1 and 4.2, it's clear that the water exchange between the Bohai Sea and  
 333 the open sea is very poor for the local freshwater budget dominate the practical salinity  
 334 magnitude and distribution. As a semi-closed sea, the seawater chemistry data of the '06's  
 335 Investigation' (Ji, 2016; Ji et al, 2021) show that nutrients and total alkalinity are significantly  
 336 different from standard seawater which is affected by the discharge of the highly sandy Yellow  
 337 River and the resuspension and dissolution of deposited calcium carbonate under the agitation of  
 338 strong winds in winter.

339 Based on the standard seawater (Millero, 2008), the total alkalinity of Bohai seawater in  
 340 winter was normalized,

$$341 \quad \Delta[\text{NTA}] = \text{TA} - 2.3 \times \frac{S_P}{35} \quad (3)$$

342 In Eq. 3, the unit of TA is  $\text{mmol} \cdot \text{kg}^{-1}$ ,  $S_P$  is the practical salinity of seawater.



345 **Figure 13.** The contour of sea surface temperature and salinity of Bohai Sea in '06's  
 346 Investigation'

347 It can be found that the  $\Delta[\text{NTA}]$  maximum value of 2.0 appears in the Bohai Bay and  
 348 Laizhou Bay on both sides of the Yellow River estuary, and its isolines spill outward in an arc  
 349 outward from the top of two bays, and the  $\Delta[\text{NTA}]$  on the bottom layer is higher than that of the  
 350 sea surface, as shown in Fig.14. A low  $\Delta[\text{NTA}]$  tongue extends westward from the Bohai Strait,  
 351 rising along its way, and its front could only reach the central of Bohai Sea which implies the  
 352 fraction of intruding Yellow Sea water continues to decline from east to the west.

353 The practical salinity is derived from the measured conductivity, temperature and  
 354 pressure of seawater by EOS-80 based on the rule of the constant relative composition of  
 355 seawater. When the relative composition of seawater changes, differences appear in practical  
 356 salinity and the mass of dissolved substances in seawater. As the seawater standard for the

357 thermodynamic properties of seawater, TEOS-10 (IOC et al, 2010) provides the equations to  
358 correct the difference of practical salinity between the mass of dissolved ionic content in  
359 seawater. The Absolute Salinity Anomaly  $\delta S_A$  in the whole Bohai Sea is always larger than 0.02  
360  $\text{g}\cdot\text{kg}^{-1}$ . The largest  $\delta S_A$  of 0.20  $\text{g}\cdot\text{kg}^{-1}$  in the Bohai Sea appears on the bottom of the Laizhou Bay  
361 in winter (Ji et al, 2021).

362 Surrounded by land on three sides, the Bohai Sea carries the sewage discharged into the  
363 sea from three provinces and one city in northern China and is under great pollution pressure. In  
364 recent years, with the rapid development of the Bohai Rim Economic Belt and the lack of  
365 reasonable planning and management, the water quality in the Bohai Sea has continued to  
366 deteriorate. worsening, red tides occur frequently (Yu et al, 2000; Shang, 2015; Xu, 2020).  
367 However, the available data are still unable to accurately calculate the magnitude of the impact  
368 of these substances on the salinity of seawater.

## 369 **5 Summery**

370 Salinity is a key variable in the modelling and observation of ocean circulation and  
371 ocean-atmosphere fluxes of heat and water. Accurate characterization of seawater salinity is a  
372 requirement for exploring the mechanisms of local ocean environmental change.

373 Based on in-situ hydrographic measurements and reanalysis data over 35 years, this paper  
374 reveals the mechanism of that the practical salinity of the Bohai Sea in January 2007 was not  
375 only significantly higher than the multi-year mean, but also the horizontal distribution was  
376 opposite to the latter, as well as indicates it is necessary to replace the Practical Salinity with the  
377 Absolute Salinity for accurately study of salinity changes of the Bohai Seawater.

378 Due to the inhomogeneity of the temporal and spatial distribution of the measurements,  
379 this paper only treats salinity variation in the Bohai Sea as a whole, the in smaller spatial scale is  
380 ignored; the temporal scale of salinity changes of this study is only accurate to the year and on  
381 seasonal or smaller time scales are not considered; moreover, it does not qualify the water  
382 exchange between the Bohai Sea and the Yellow Sea. At present, a study on the characteristics  
383 and patterns of smaller temporal scale salinity changes in the Bohai Sea, combined with ocean  
384 models and other measurement data, is currently underway.

385 Due to material transport by runoff, re-dissolution and mixing of bottom materials,  
386 biological uptake and decomposition, relative composition of offshore seawater may be  
387 significantly different from that of the open ocean, and seawater salinity is no longer a  
388 conservative quantity. However, the available information cannot precisely calculate the  
389 magnitude and distribution change pattern of the impact of these dissolved material anomalies on  
390 the salinity of the Bohai Sea.

391

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