

Elucidating geochemical heterogeneity and evolution of the explosively erupted Curacautín magma, Llaima volcano, Chile

Jade Bowers¹, V. Dorsey Wanless¹, Darin Schwartz¹, Brittany Brand¹, and Benjamin Andrews²

¹Boise State University

²National Museum of Natural History, Smithsonian Institution

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Abstract

Explosive mafic eruptions are among Earth's most hazardous volcanic phenomena due to the rapid ascent rates of mafic magma limiting time for early detection and warning systems. To date, work on explosive mafic eruptions has primarily been in the context of conduit and mafic magma fragmentation processes. Our work uses geochemistry and numerical modeling to establish magma source and system dynamics that led to an explosive mafic eruption at Llaima volcano, Chile. This study compares major and trace elements of 70 whole rock ignimbrite samples collected from 17 different outcrops of the Curacautín ignimbrite, including four extensive outcrops sampled vertically every 1-3 m to examine chemical variations and changes in magma source prior to and during the eruption. The bases and tops of the four stratigraphic sections and four additional samples with unique chemical signatures were analyzed for Sr-Nd-Pb isotopes to investigate mantle source variation. Eruptive products of the Curacautín eruption define a medium-K calc-alkaline suite of basalt to andesite (SiO₂ 50-58 wt. %). Both regionally and stratigraphically, trace element patterns produced remarkably similar, parallel patterns with the most enriched sample in the west where the base of the ignimbrite is exposed ($La_N/Sm_N = 1.63$). The Curacautín ignimbrite is chemically heterogeneous [e.g., MgO=2.5-6 wt. %, Ce = 11-39 ppm, and Ba/Th= 170-263], but Pb isotopes are homogeneous ($^{207}Pb/^{204}Pb = 15.59-15.61$), indicating the magma source was influenced by slab inputs or secondary process (e.g., AFC). However, Sr-Nd isotopes still need to be collected. To constrain mantle source heterogeneity, we utilize hierarchical clustering analysis (HCA). HCA results indicate four chemically distinct groups in the north, two in the east, and two in the west. To explore secondary processes, Magma Chamber Simulator (MCS; Bohron et al., 2014) modeling will be used to produce thermodynamic models to constrain the chemical variability resulting from fractional crystallization (FC) processes.



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Jade M. Bowers^{1*}, V. Dorsey Wanless¹, Darin M. Schwartz¹, Brittany D. Brand¹, Ben J. Andrews²

¹Department of Geosciences, Boise State University, Boise, ID 83713, USA

²Global Volcanism Program, National Museum of Natural History, Smithsonian Institution, Washington, DC 20560, USA

Introduction & Background

1. Problem

Mafic explosive eruptions are among Earth's most hazardous volcanic phenomena due to their rapid magma ascent rates limiting the time frame for eruption early detection and public warning systems [ref. 1-6]. Despite this concern, we do not understand how these eruptions geochemically differ from other mafic eruptions at the same volcanic center.

2. Approach

Here we present a case study focusing on the petrogenesis of the Curacautín ignimbrite, a mafic ignimbrite at Llaima Volcano, Chile, produced by a mafic explosive eruption around 12.6 ka [ref. 7]. Geochemical variability is investigated using major and trace element ratios, radiogenic isotopes, and agglomerative hierarchical clustering analysis to distinguish between fractional crystallization and source material heterogeneity.

3. Curacautín Ignimbrite Geochemistry

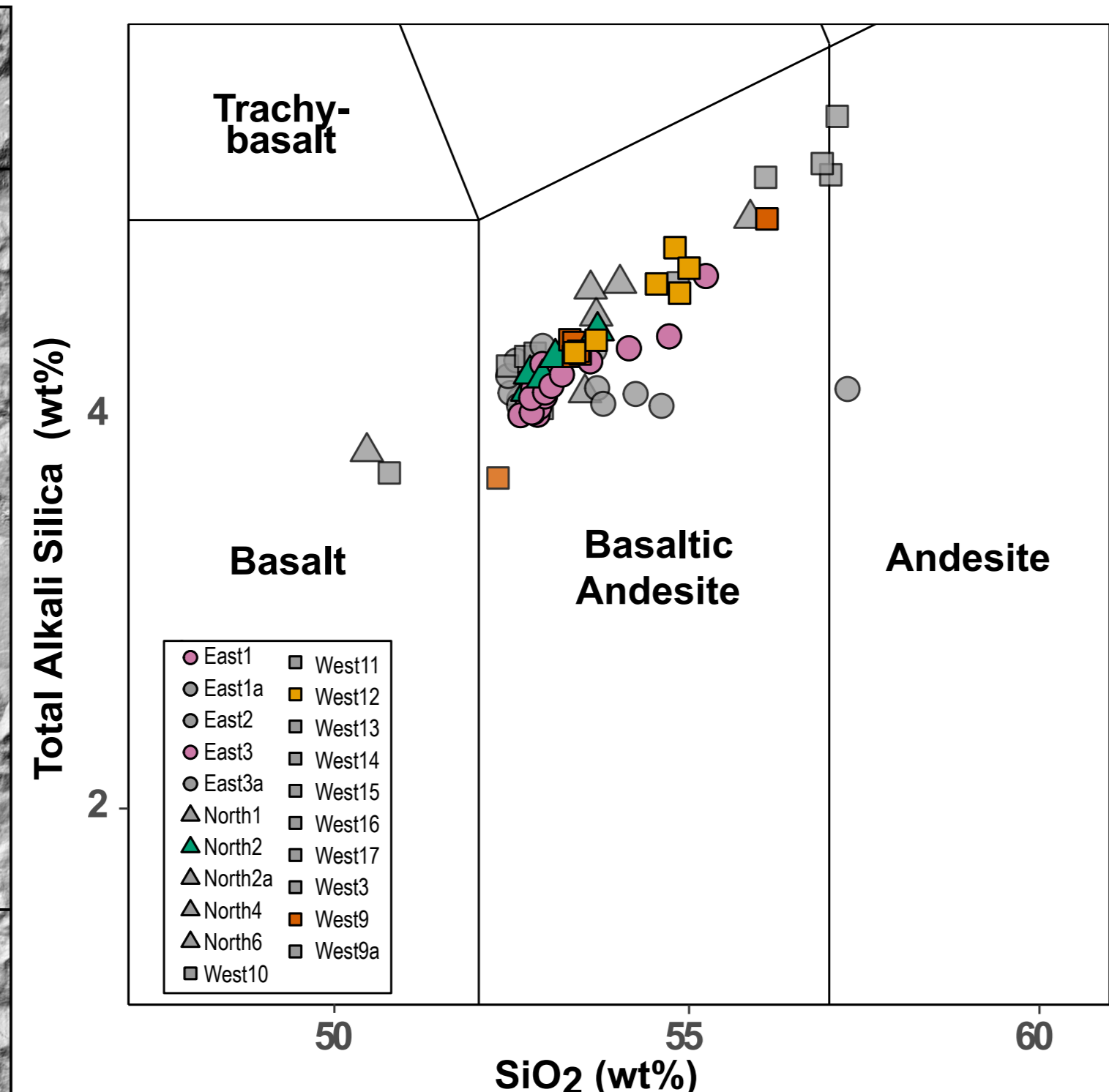
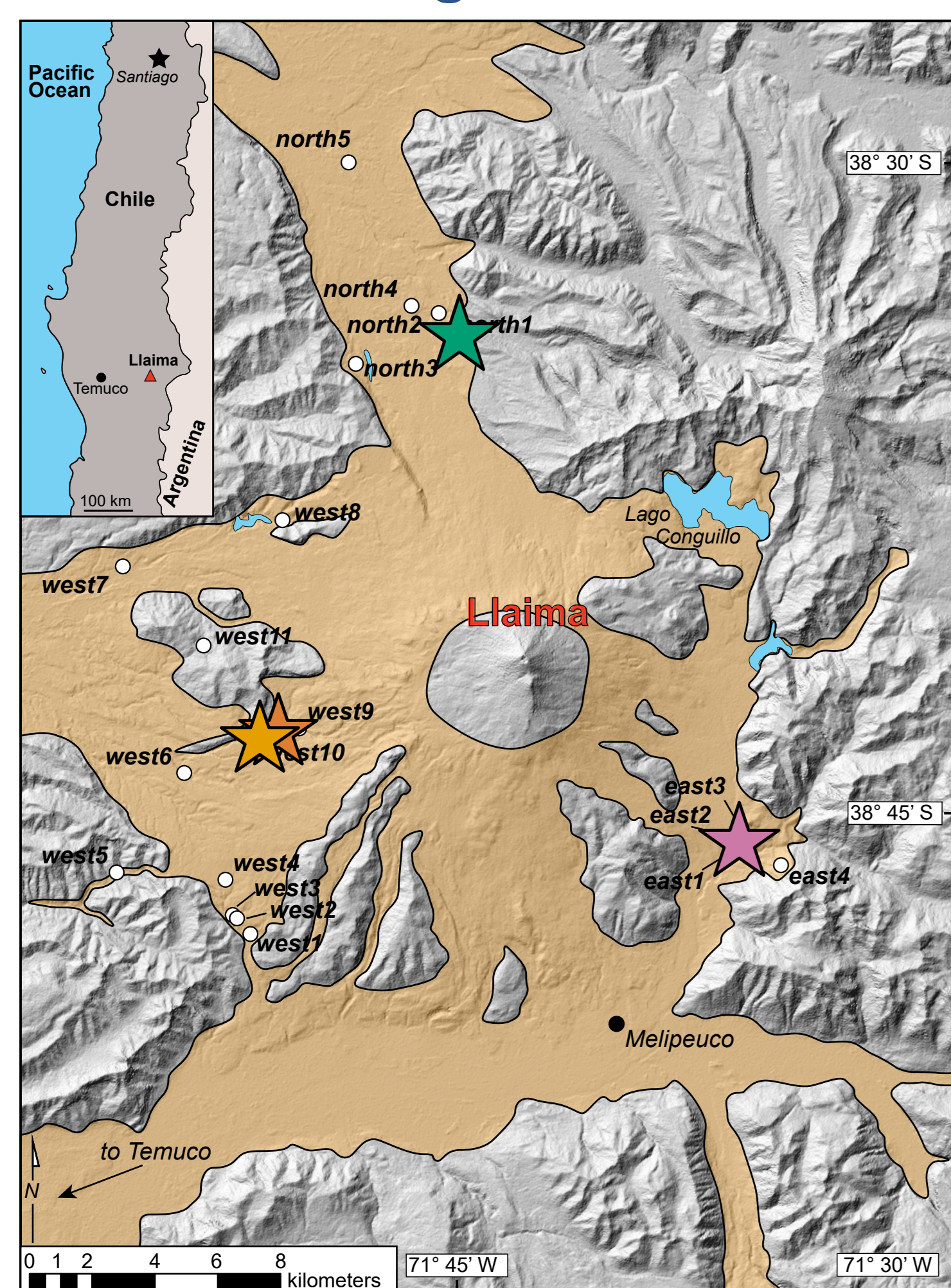


Figure 1. Map showing the location of Llaima volcano, Chile (left). Yellow polygon represents mapped extent of Curacautín ignimbrite. Locations of sampled outcrops are indicated by white circles and the four major exposures are indicated by stars. Total Alkali versus Silica [TAS] (top). Samples collected from the four key exposures are highlighted, all other samples are gray. Curacautín is primarily basaltic andesite.

1. Hierarchical Clustering

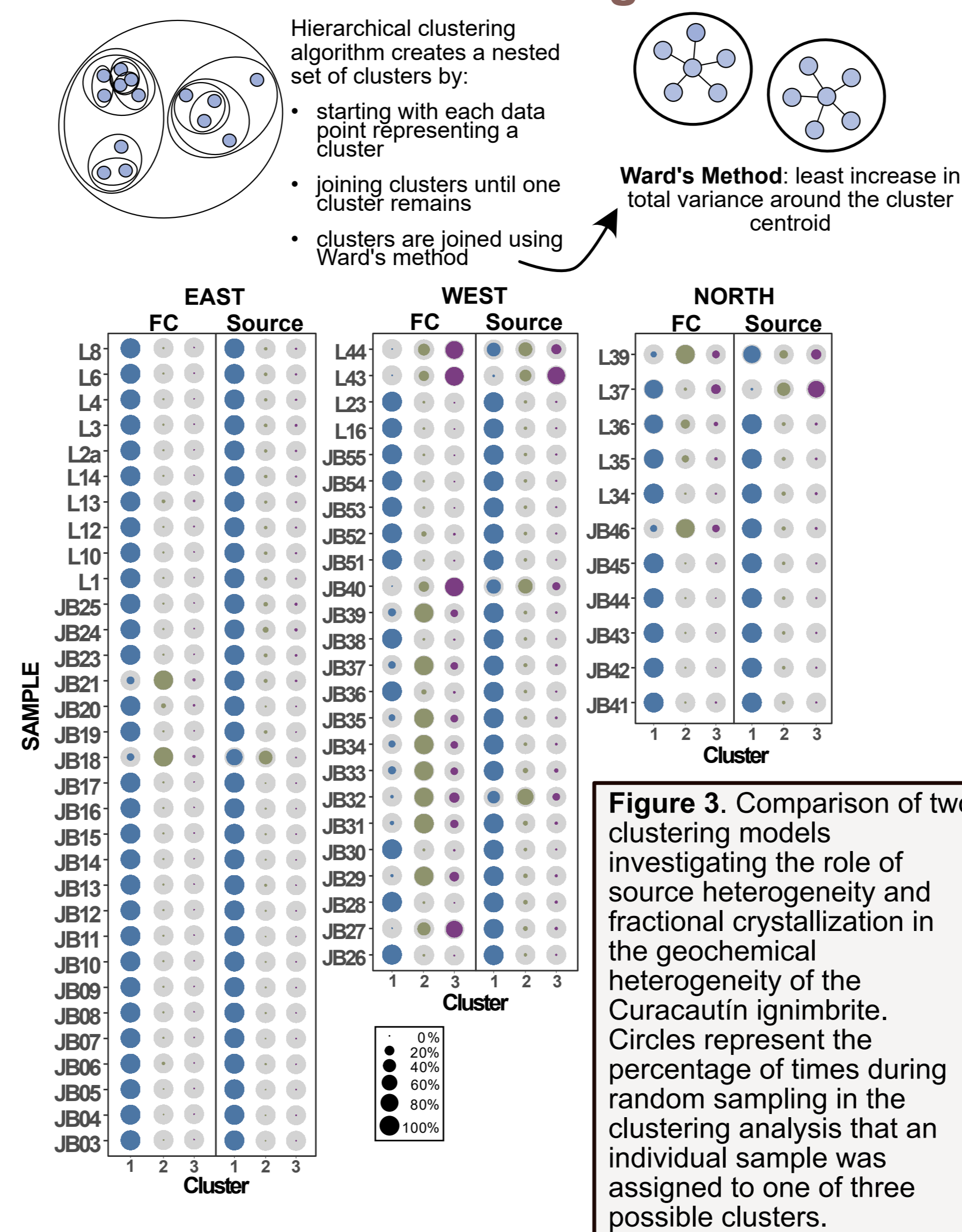


Figure 3. Comparison of two clustering models investigating the role of source heterogeneity and fractional crystallization in the geochemical heterogeneity of the Curacautín ignimbrite. Circles represent the percentage of times during random sampling in the clustering analysis that an individual sample was assigned to one of three possible clusters.

- Unratified REE concentrations are used in the FC model because concentrations will vary as fractionation occurs
- REE are ratioed to remove effect of FC on concentrations and focus on subduction processes influence on source material compositions
- Fewer clusters are stable in the Source model
- More clusters in the FC model indicates FC is the dominant process influencing geochemical variability
- Samples that are assigned cluster 2 or 3 in Source model indicate potential mixing end members

3. Subducted sediments

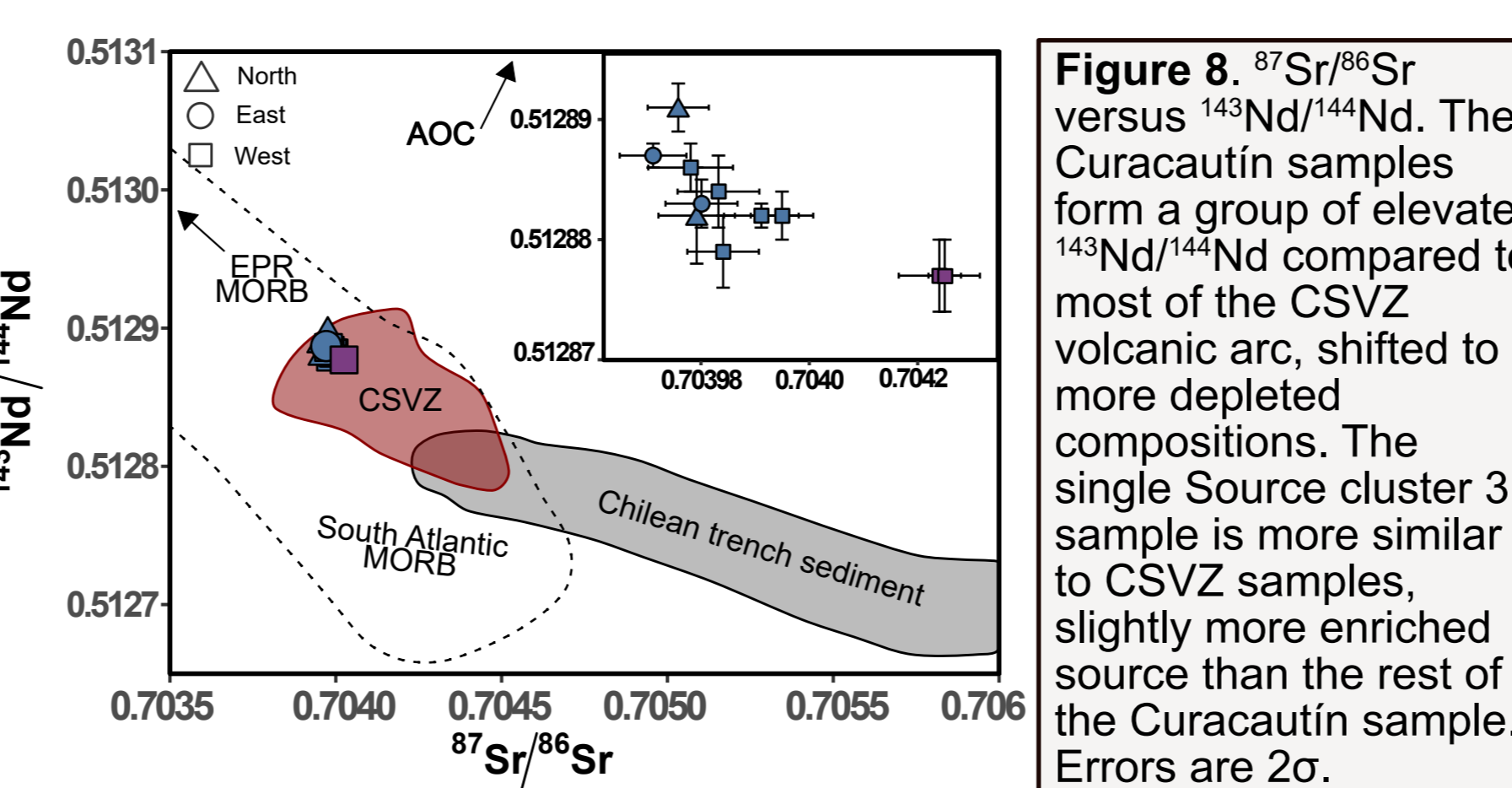


Figure 8. ⁸⁷Sr/⁸⁶Sr versus ¹⁴³Nd/¹⁴⁴Nd. The Curacautín samples form a group of elevated ¹⁴³Nd/¹⁴⁴Nd compared to most of the CSVZ volcanic arc, shifted to more depleted compositions. The single Source cluster 3 sample is more similar to CSVZ samples, slightly more enriched source than the rest of the Curacautín sample. Errors are 2σ.

4. Fluid Flux and Degree of Melting

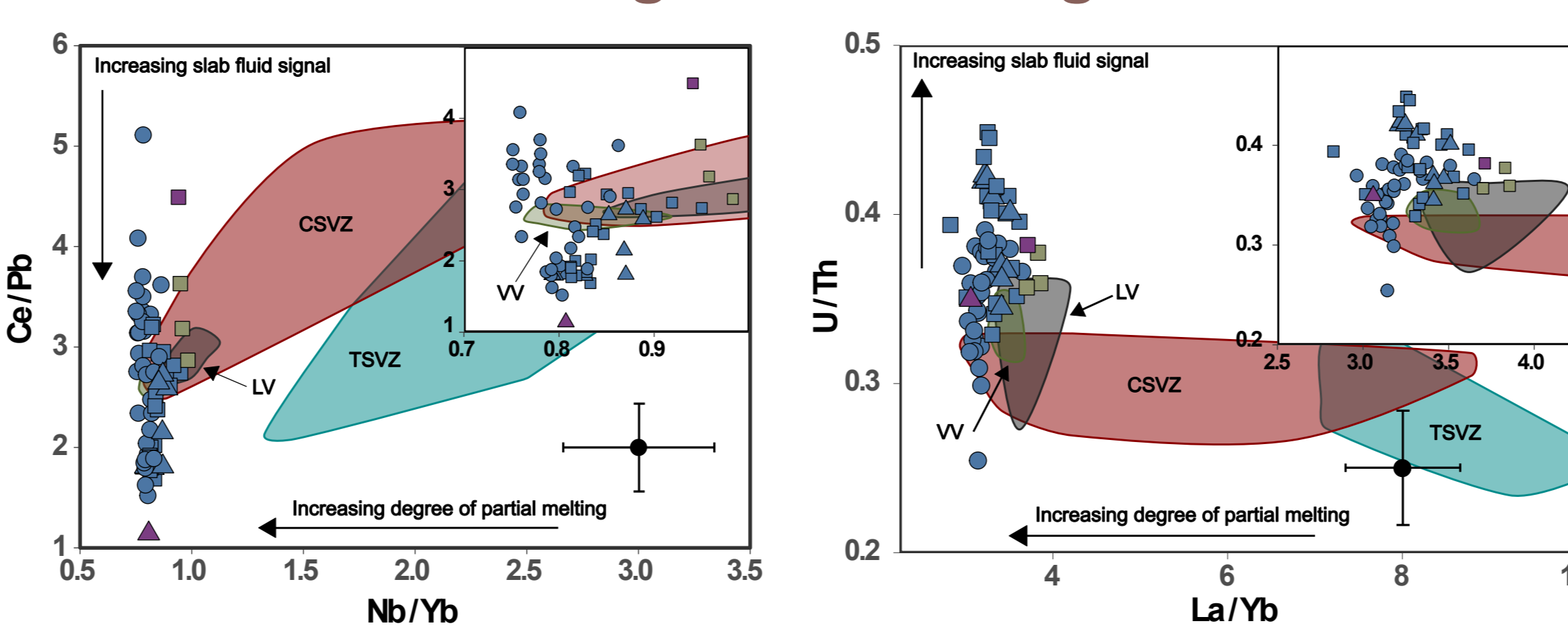


Figure 9. Nb/Yb versus Ce/Pb (left). The Curacautín samples define a vertical trend with variation in Ce/Pb, potentially reflecting mixing between a high slab fluid source and an end member with less slab fluid influence, and the lowest Nb/Yb of the CSVZ (red), TSVZ (blue), Villarrica (green), and other published Llaima (gray) data, reflecting the highest degrees of melting or derivation from a more depleted source material. La/Yb versus U/Th (right). The Curacautín samples show some vertical spread in U/Th and overlap with the lowest La/Yb values in the CSVZ volcanic arc consistent with a higher slab component and higher degree of melting/more depleted source material. Errors are 2σ.

2. Assimilation and Fractional Crystallization

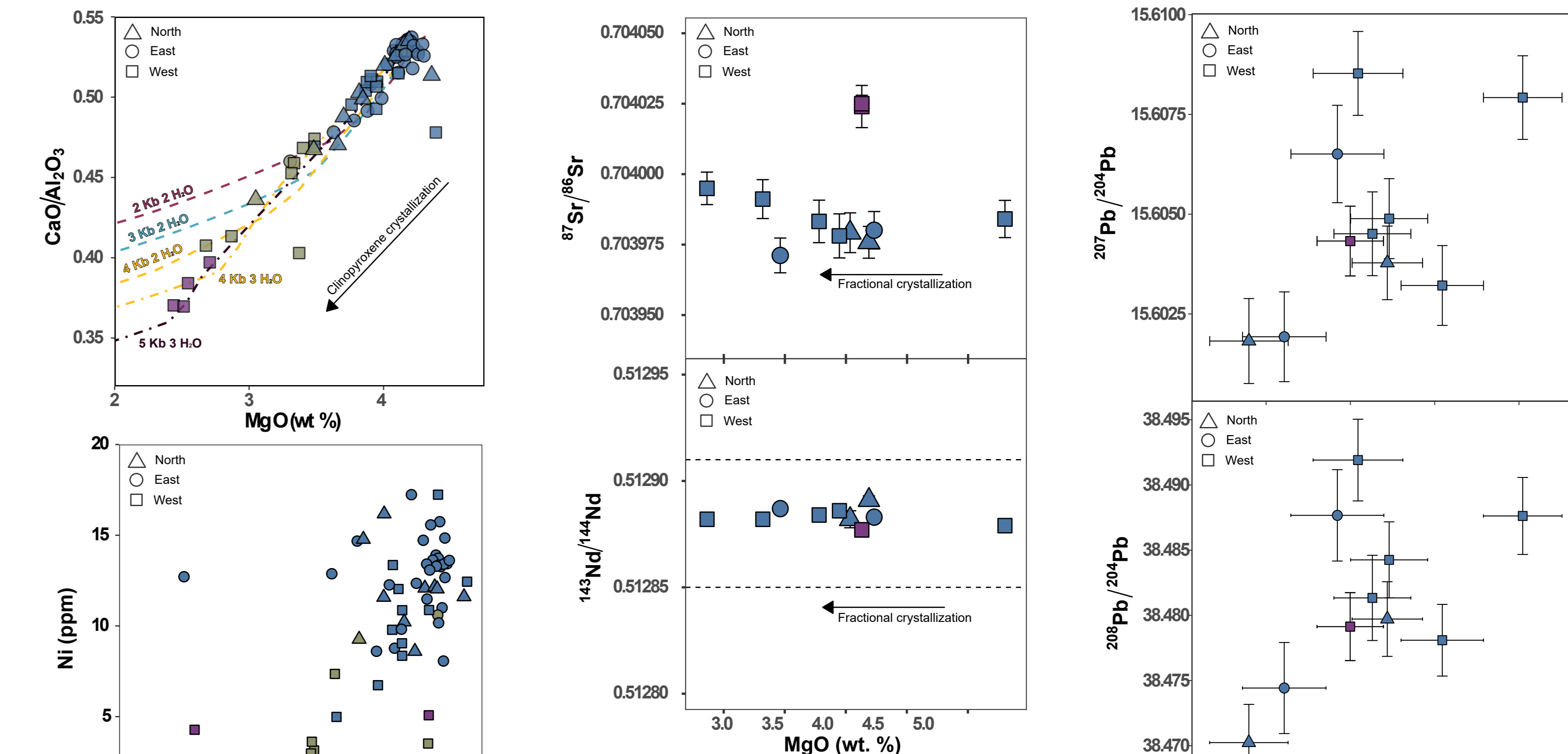


Figure 4. (top) Cluster assignments from the FC model are plotted with MCS fractionation paths to highlight FC as the dominant process influencing the geochemical variability of the samples. Colors match those assigned to the three clusters in Figure 3. The FC model cluster 1 (blue) contains the least fractionated of the samples and FC model cluster 3 (purple) is primarily the most fractionated. (bottom) Cluster assignments and compatible trace elements (e.g. Ni in Olivine) highlight multiple evolution trajectories.

Figure 7. MgO versus Ba/Th (right). Vertical trend in Ba/Th with little variability in MgO (wt. %) at 4.1 wt. % reflecting mixing with L43, potentially enriched in Ba by assimilation of young crustal roots. Errors are 2σ.

Figure 5. MgO versus (top) ⁸⁷Sr/⁸⁶Sr, (bottom) ¹⁴³Nd/¹⁴⁴Nd for Curacautín ignimbrite. The data can be explained by crystal fractionation indicated by the horizontal arrow. The only suggestion of contaminant assimilation is the deviation in ⁸⁷Sr/⁸⁶Sr between cluster 1 samples and the cluster 3 sample. Errors are 2σ.

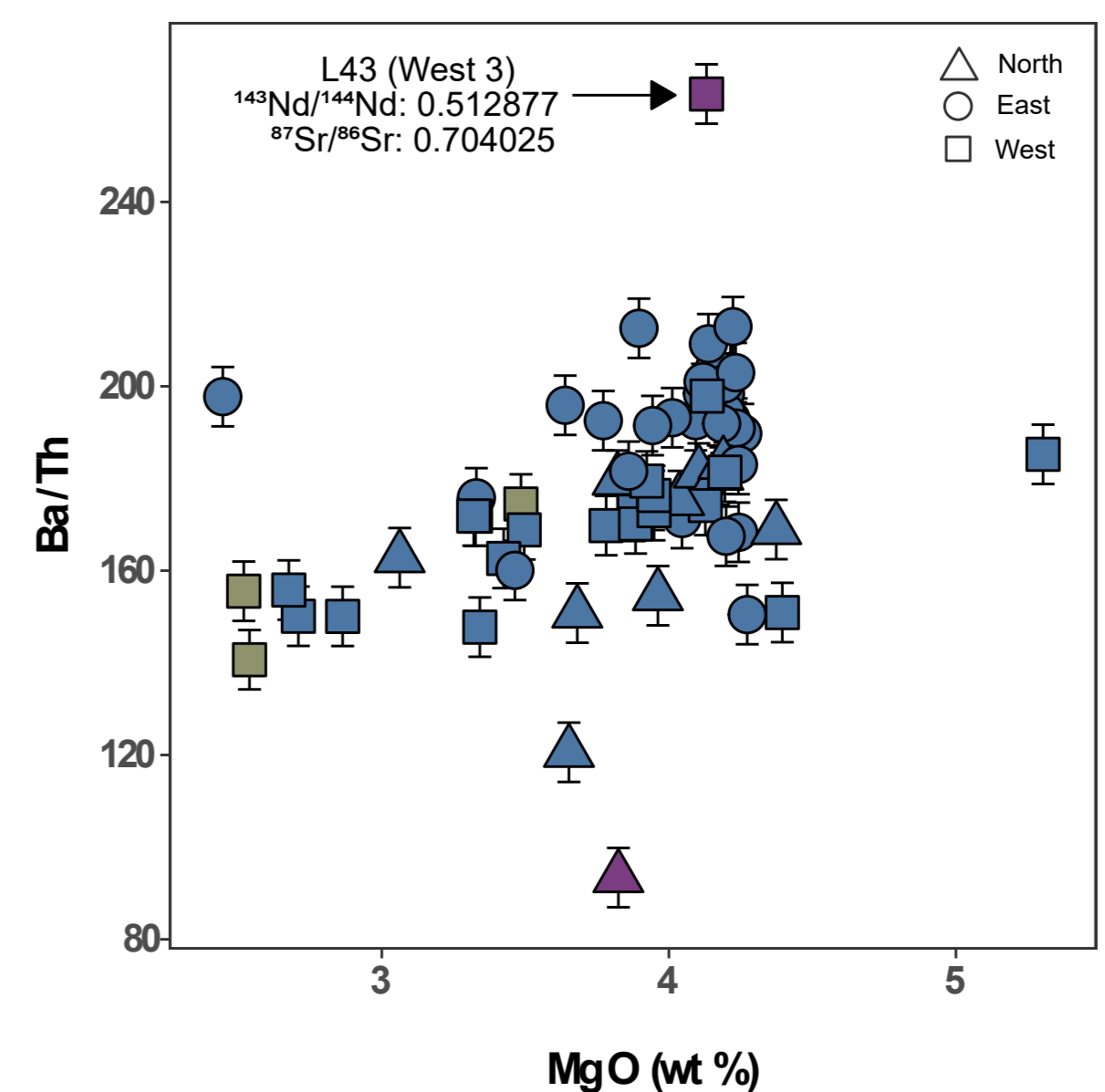


Figure 6. ²⁰⁶Pb/²⁰⁴Pb versus (top) ²⁰⁷Pb/²⁰⁴Pb, (middle) ²⁰⁸Pb/²⁰⁴Pb, and (bottom) ⁸⁷Sr/⁸⁶Sr. Source clustering model does not explain variability in ²⁰⁷Pb/²⁰⁴Pb, ²⁰⁸Pb/²⁰⁴Pb, or ²⁰⁶Pb/²⁰⁴Pb, but does capture variability in ⁸⁷Sr/⁸⁶Sr. Errors are 2σ.

Interpretations

- MELTS fractional crystallization models capture CPX crystallization at 5000 bars with 3 wt. % H₂O in the melt
- L43 (West 3) slightly elevated ⁸⁷Sr/⁸⁶Sr in comparison to other Curacautín samples and high Ba/Th reflecting an mixing end member representing assimilation of young crustal roots
- No evidence for sediment derived melt in Curacautín mantle source
- High slab-derived fluid flux can be explained by subduction of the more hydrated crust in the Valdivia Fracture Zone
- High fluid flux causes elevated degree of partial melting in the mantle wedge
- Fractional crystallization is the primary process driving geochemical variability
- A secondary mixing signal is highlighted in the clustering analysis, MgO versus selected trace elements, and radiogenic strontium
- We propose assimilation of young crustal roots and high slab-derived fluid flux in the mantle wedge influenced the melt source material of the Curacautín magma.

The signature of slab-derived fluid flux (e.g., Ce/Pb, U/Th) of the Curacautín samples is higher than any other CSVZ deposits, potentially influencing eruption dynamics of the Curacautín magma.

Acknowledgments References

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¹Coltelli et al., 1998. *Geology*. ²Sable et al., 2006. *JVGR*. ³Szramek et al., 2006. *JVGR*. ⁴Szramek, 2016. *JGR: Solid Earth*. ⁵Arzilli et al., 2019. *Nature Geoscience*. ⁶Valdivia et al., 2021. *BullVolc.* ⁷Marshall et al., 2021. *JVGR*.